

GARNETS OF MICA-SCHISTS FROM CRNI VRH AND RESAVSKI HUMOVI (SERBIA)

S.Erić, D.Milovanović and D.Babič

Department of Mineralogy and Petrology, Faculty of Mining and Geology, Belgrade University, Serbia and Montenegro, suzanaeric@yahoo.com

ABSTRACT

Metamorphic rocks of the Crni Vrh and Resavski Humovi area belong to the north-northwestern part of the Serbo-Macedonian Composite Terrane and consist of different types of mica-schists, gneisses and amphibolites.

Mica-schists are composed of biotite, muscovite, garnet, staurolite, kyanite, small quantity of K-feldspar, acid plagioclase and quartz. Accessory minerals are apatite, zircon, tourmaline and ilmenite.

Garnets in mica-schists from both localities occur as subhedral to anhedral porphyroblasts from 0.2 to 10mm in size. Most garnets are oriented parallel to the dominant foliation.

Garnets from the Crni Vrh mica-schists contain more spessartine component than garnets from Resavski Humovi, while almandine component show opposite tendency. The difference between those garnets is clearly visible, as well as zonality expressed through the increase of almandine component from center to the grain rim. The spessartine component as mentioned above, shows the opposite tendency.

Using different geothermometers and geobarometers it was found that mica-schists from Crni Vrh were formed at temperatures from 550-600°C and pressures from 6-6.2 kbar, while mica-schists from Resavski Humovi developed at slightly lower temperatures (520-560°C) and at some higher pressures (7-7.3 kbar).

1 INTRODUCTION

Garnet is an important mineral in the study of metamorphic processes because individual crystals preserve a record of the metamorphic and deformational environment of the rock during one or more tectonic events. Metamorphosed metapelites also provide useful information, as pelitic bulk composition is generally mineralogical sensitive to changes in pressure and temperature. Investigated mica-schists with garnets belong to metamorphic rocks of the Serbo-Macedonian Composite Terrane (SMMT) (Karamata and Krstić, 1996) which are composed of different types of gneisses, mica-schists, amphibolites and rarely marbles. Migmatites and rare eclogites occur along the eastern margin of this unit.

Metamorphic rocks of the SMMT show nearly ubiquitous amphibolite-facies metamorphic assemblage and are interpreted as products of prograde Variscan metamorphism.

K/Ar radiometric age determinations on hornblende (270 ± 10 Ma) indicate that the major metamorphic phase occurred during Variscan age, whereas the K/Ar ages of biotite, (120 ± 10 Ma), muscovite and feldspar are probably related to Alpine tectonic events.

2 GEOLOGY

Metamorphic rocks of the Crni Vrh and Resavski Humovi area belong to the north-northwestern part of the SMMT (Fig. 1) and consist of different types of mica-schists, gneisses, and amphibolites. These rocks represent a metamorphosed flysch-like "series", cut by various igneous rocks.

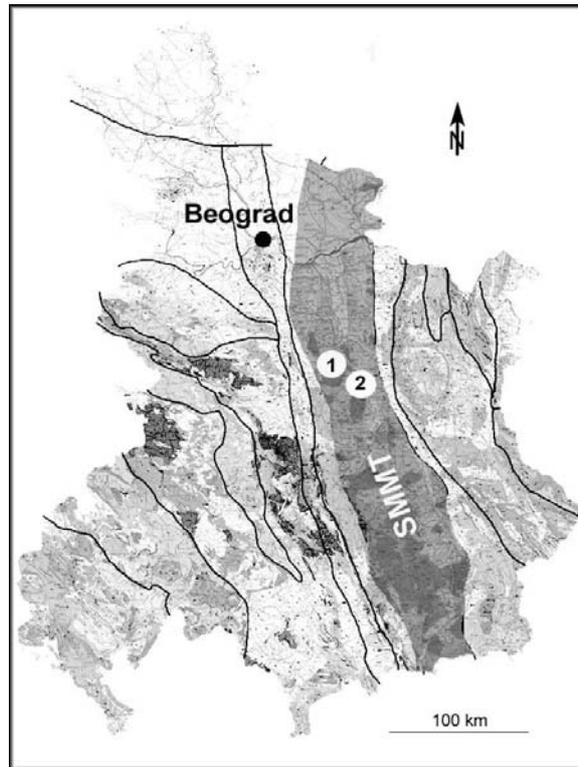


Figure 1. Position of mica-schists from the Crni Vrh (1) and Resavski Humovi (2) in Serbo-Macedonian Composite Terrane (Karamata and Krstić, 1996)

Mica-schists of the Crni Vrh occur 10km north of Bagrdan on the road between the villages Strižilo and Donja Komarica, along Pišljak and Kusača streams. Mica-schists of the Resavski humovi are located 15 km southwest of Svilajnac and 2km of Miljkov monestery.

These rocks are medium to coarse grained, showing migmatitic features such as stromatic and stictolithic structures with rare intercalations of amphibolite, marble and quartzite. Isolated rootless, mostly isoclinal folds of cm-scale quartz veins and isoclinal folds in quartzitic or feldspar-rich layers are parallel with the main schistosity and point an early deformation event. The mica-schists sometimes gradually pass into to gneisses and migmatites. When abundance and size of the K-feldspar porphyroblasts increase feldspatic to quartzo-feldspatic streams occur. These rocks show a pervasive high grade S_1 foliation that is overprinted by a slightly to strongly discordant S_2 foliation. Locally crenulation cleavage grades into a continuous schistosity by a process of progressive transposition and recrystallization.

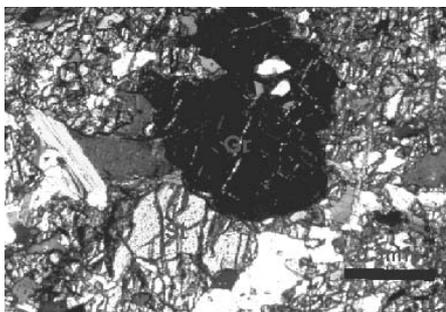


Figure 2. Porphyroblasts of garnet and staurolite



Figure 3. Porphyroblasts of garnet with micas

3 PETROGRAPHY

Mica-schists are composed of biotite, muscovite, garnet, staurolite, kyanite, small quantity of K-feldspar, (orthoclase and microcline), acid plagioclase, and quartz. Accessory minerals are apatite, zircon, tourmaline and sphene.

The texture of mica-schists is lepidoblastic, granoblastic and porphyroblastic with variable size of minerals, from a few mm to a few cm. Biotite and muscovite make up to 50% vol of these rocks (Table 1).

Biotite and muscovite vary in lengths from 0.1 to 2 cm and show a strong shape preferred orientation defining schistosity in mica-schists. Three generations of micas are visible: deformed flakes up to 2 cm in size, which are most probably relics of the clasts protolith; a second generation, in the matrix, is fine to medium grained, synkinematic, and finally youngest postkinematic generation are represented by fine flakes randomly oriented, which cut the earlier generations. An anastomosing configuration of biotite and muscovite around large garnet porphyroblasts indicates that the formation of the dominant schistosity postdated garnet growth.

Biotite and muscovite also occur as inclusions in garnet and staurolite porphyroblasts. Biotite is usually replaced with chlorite and fine grained Fe-oxides.

Table 1. Mineral assemblages and modal composition of the mica-schists from the Crni Vrh and Resavski Humovi

Crni Vrh	Resavski Humovi
Qz(13%) + Ms(44%) + Bi(11%) + Gt(18%) + Pl(3%) + Kf + St(10%) + Tour + Chl + Ru + Ilm	Qz(20%) + Ms(41%) + Bi(13%) + Gt(16%) + St(5%) + Ky(2%) + Pl(2%) + Kf + Tour + Ru + Ilm

Garnet shows various textures. It occurs as subhedral to anhedral porphyroblasts (Fig. 2; Fig. 3) from 0.2 to 10 mm in size, sometimes partially replaced by aggregates of chlorite and biotite. Garnet usually contains sigmoidal to spiral inclusion of quartz, micas and ilmenite, suggesting a synkinematic crystallization or intertectonic growth over preexisting crenulation cleavage. Some garnets contain inclusions only in their interior part. Most garnets are oriented parallel to the dominant foliation. Small crystals of garnet occur in strings along foliation planes in many samples which suggests a pre to syn-kinematic crystallization. Some garnets are undeformed and clearly post-date penetrative deformation.

Staurolite in mica-schist from the Crni Vrh occurs in impressive size porphyroblasts (up to 10 cm in diameter) and contain inclusions of garnet, micas (mainly biotite) quartz and ilmenite which locally comprise 10% or more of the porphyroblasts.

Quartz exhibit undulatory extinction and occurs as an interpenetrative mosaic of equal grains, up to 1 mm in size that define ribbons parallel to foliation. This mineral is abundant both as matrix and inclusions in garnet. In the ribbons quartz sometimes displays coarser grain size, straight grain boundaries and triple junctions that indicate grain growth processes driven by reaction of interfacial free energy. Quartz also occurs as recrystallized grains in the pressure shadows around grains of garnets.

K-feldspar is rare, occurring either as subhedral or anhedral grains in the matrix, usually associated with quartz.

4 CHEMISTRY

The petrographic study of these metamorphic rocks has been supplemented by electron microprobe analysis of their main minerals. Major element analyses were obtained for garnet, staurolite, biotite, muscovite and K-feldspar in the Laboratory of Mineralogy in Hamburg (Germany) and in the Laboratory of Mineralogy, at the Aristotle University of Thessaloniki (Greece). Representative compositional data are available from the authors on request.

About 50 mineral analyses of garnets were carried out and structural formulas were calculated.

Garnets from the Crni Vrh mica-schists contain more spessartine component than garnets from the Resavski humovi, while almandine component shows the opposite tendency. Difference between these garnets is clearly visible in Fig. 4, as well as zonality expressed through the increase of

almandine component from the center to the grain rim. The spessartine component as mentioned above, shows the opposite tendency.

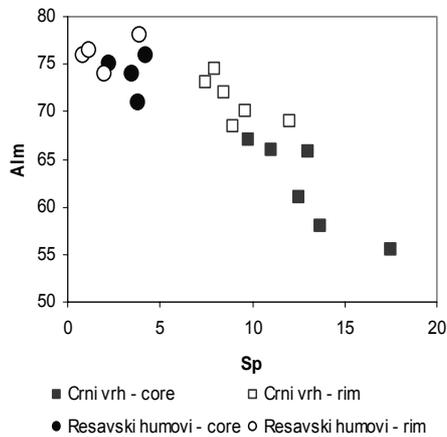


Figure 4. Relation of almandine and spessartine component in garnets

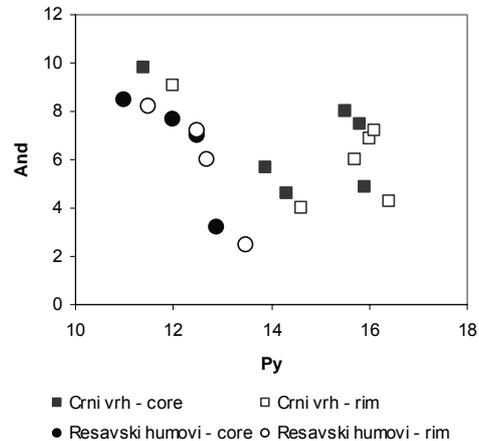


Figure 5. Relation of pyrope and andradite component in garnets

5 X-RAY ANALYSIS

Concerning the chemistry of the analyzed garnets higher values for unit cell parameters (Table 2) have been expected for garnets in mica-schists from Crni Vrh (with more spessartine and lower almandine content) than for those in mica-schists from the Resavski Humovi. However, a completely opposite trend was noted. According to well known fact that garnet unit cell increases from pyrope toward andradite, this anomaly should be attributed to higher pyrope (Fig. 5) and lower andradite content in the former than in the latter one.

Table 2. Unit cell parameters of garnet from investigated mica-schists

Unit cell parameters of garnets from Crni Vrh	Unit cell parameters of garnets from Resavski Humovi
$a_o = 11.558 (2) \text{ \AA}; V_o = 1543.9 (8) \text{ \AA}^3$	$a_o = 11.571 (2) \text{ \AA}; V_o = 1549.3 (7) \text{ \AA}^3$
$a_o = 11.552 (2) \text{ \AA}; V_o = 1541.6 (7) \text{ \AA}^3$	$a_o = 11.568 (2) \text{ \AA}; V_o = 1547.9 (8) \text{ \AA}^3$
	$a_o = 11.5823 (8) \text{ \AA}; V_o = 1553.7 (3) \text{ \AA}^3$

6 GEOTHERMOMETRY AND GEOBAROMETRY

Samples for thermobarometry were selected on the basis of petrographic criteria suggesting equilibrium in millimeter scale. Mineral assemblages were compared with petrogenetic grid and chemical compositions of minerals and were used to calculate temperature and pressure using published calibrations of exchange equilibria. Several geothermometers and geobarometers have been used to determine P-T conditions (Fig. 6) during formation and evolution of the investigated mica-schists (Table 3).

Table 3. Calculated temperatures and pressures of the metamorphism of the mica-schists from Crni Vrh and Resavski Humovi on basis of different methods

Methods	Mica-schists from Crni Vrh	Mica-schists from Resavski Humovi
	T (°C)	T (°C)
Ferry and Spear, 1978	577-581	518-528
Hodges and Spear, 1982	603-625	554-564
Perchuk and Lavrentova, 1983	599-610	564-574
Ganguly and Saxena, 1984	583-598	527-567
Indares and Martignole 1985	559-596	512-522
Hynes and Forest, 1988	590-602	544-559
Gessman et al., 1997	521-539	486-502
	P (kbar)	P (kbar)
Powell and Holand, 1988	6.0-6.2	7.0-7.3

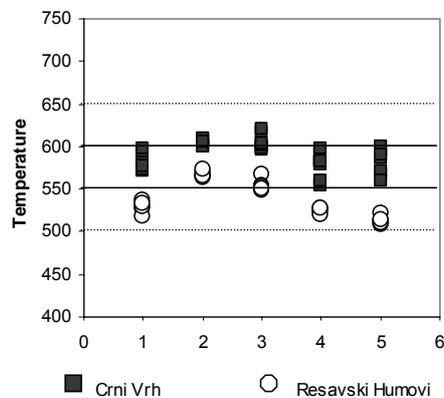


Figure 6. Calculated temperatures of metamorphism of the mica-schists from the Crni Vrh and Resavski Humovi

1-Ferry and Spear, 1978, 2-Perchuk and Lavrentova, 1983, 3-Hodges and Spear, 1982, 4-Ganguly and Saxena, 1984, 5-Indares and Martignole 1985

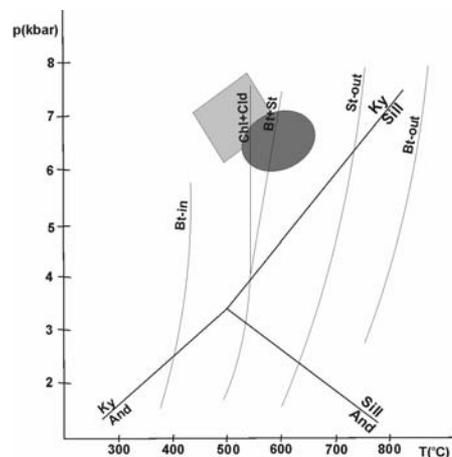


Figure 7. Position of metamorphic assemblages from mica-schist from Crni Vrh (dark gray field) and Resavski Humovi (light gray field)

7 CONCLUSION

Textural relationships and petrogenetically significant mineral assemblages in the investigated rocks imply peak P-T conditions of 6-7 kbar and temperatures ranging from 550-600°C. Calculated P-T conditions agree with the metamorphic conditions and the P-T path deduced through phase equilibria in pelitic systems (Fig. 7).

The protoliths of mica-schists were the equivalent of a sedimentary unit consisting of mudstone, carbonaceous shale, greywackes, perhaps deposited in epicontinental or intracontinental basin.

Metamorphic assemblages and field observations suggest a complex tectonic history. We identify three generations of minerals based on mineral paragenesis, mineral-fabric relationships and disequilibrium assemblages.

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